

# SNOW AND ICE FAR-INFRARED SPECTRAL EMISSIVITY RETRIEVALS FROM FORUM-LIKE MEASUREMENTS

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- In high-elevation locations and polar regions where the dry and cold atmosphere is not opaque in the far-infrared (FIR), part of the surface emission can escape to space
- In the FIR, snow surface emissivity shows considerable spectral variation
- The impact of modelled FIR surface emissivity was studied by Chen et al. 2014 and Feldman et al. 2014; Huang et al. 2016 compiled a database of surface spectral emissivity
- There is only a handful of retrievals of infrared spectral surface emissivity reaching the far-infrared
- Bellisario et al. 2016 and Murray et al. 2020 exploited observations taken by TAFTS during an airborne campaign over Greenland

## ESA-DEVELOPED EARTH OBSERVATION MISSIONS



## Science



## Copernicus

## Meteorology



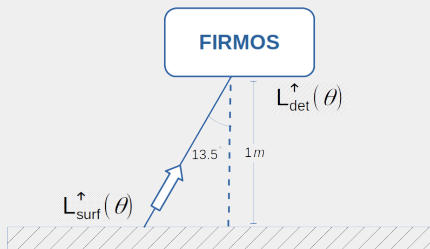
- FTS, 100–1600  $\text{cm}^{-1}$  (6.25–100  $\mu\text{m}$ ), spectral resolution is 0.5  $\text{cm}^{-1}$
- One of FORUM objectives is the retrieval of FIR surface emissivity
- Excellent sampling at high latitudes (sun-synch 98.7° orbit)
- Uncertainty of 0.01 in 300–600  $\text{cm}^{-1}$ , 50  $\text{cm}^{-1}$



- FORUM demonstrator
- FTS, 100–1000  $\text{cm}^{-1}$ ,  
 $\Delta\nu = 0.3 \text{ cm}^{-1}$  ground-based,  
balloon
- Zugspitze campaign (German  
Alps, 3000 a.m.s.l)

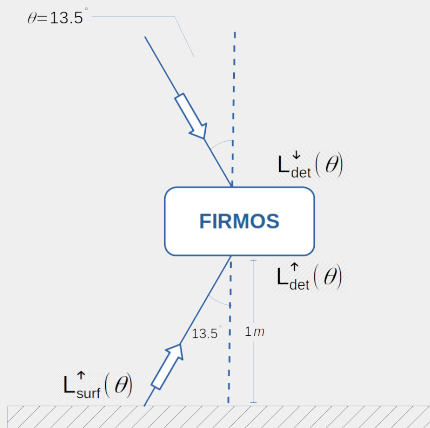
# FIRMOS

- FORUM demonstrator
- FTS,  $100\text{--}1000\text{ cm}^{-1}$ ,  
 $\Delta\nu = 0.3\text{ cm}^{-1}$  ground-based,  
balloon
- Zugspitze campaign (German  
Alps, 3000 a.m.s.l)
- Upwelling radiance @13.5  
degree off-nadir

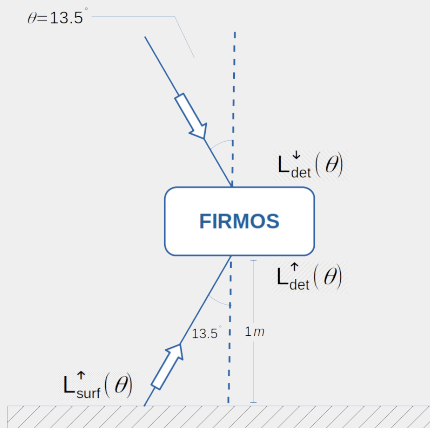


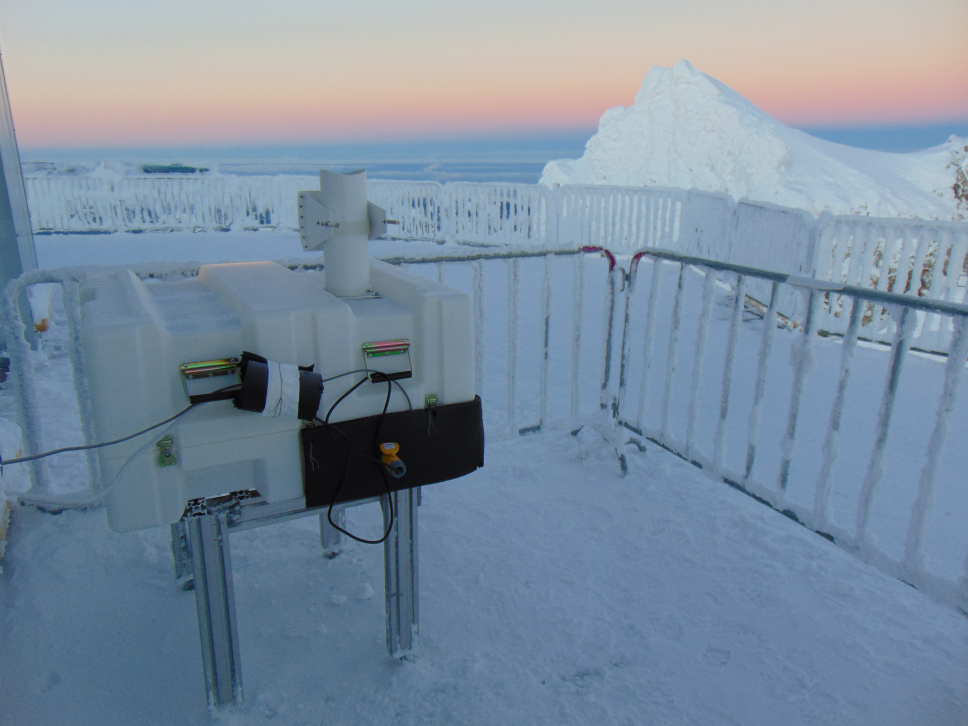
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degree viewing zenith angle



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- Downwelling radiance @13.5  
degree viewing zenith angle
- 9 samples characterised by  
snow grain type, density  
( $\text{kgm}^{-3}$ ), and specific surface  
area ( $\text{SSA}, \text{m}^2\text{kg}^{-1}$ )



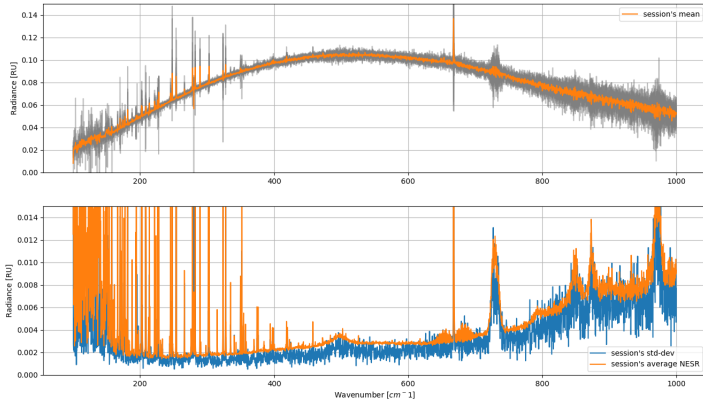




# SNOW CHARACTERISATION

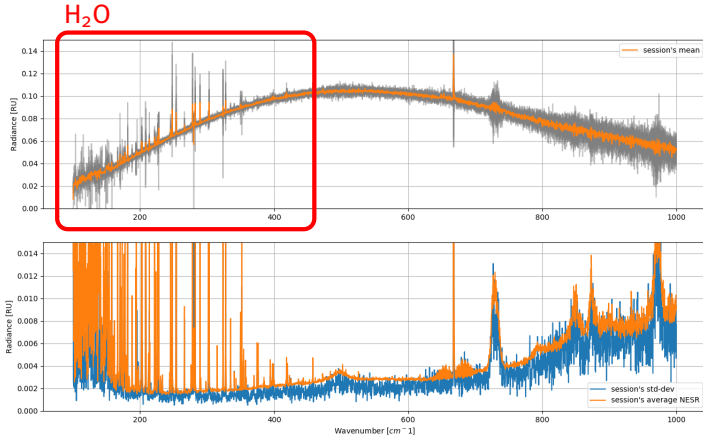
id	name	description	spectra	mins
17:30 - d1	SKY		1*	8
17:51	ICE	Ice	10	79
19:38	FS	Fresh snow	12	95
21:36	ACC30	Accumulated 30cm snow	11	87
23:42	SKY		1*	8
06:19 - d2	SKY		1*	8
06:44	DH	Depth hoar snow	13	103
09:01	DHRUG	Depth hoar snow with grooves	12	95
10:52	RGREF	Melted forms snow	7	56
12:13	SKY		1*	8
13:00	MFREF	Refrozen snow - natural case	14	111
15:00	SKY		1*	8
16:20	SKY		1*	8
16:31	DS	Dense slab snow	14	111
18:48	ACC100	Accumulated (100 cm) snow	13	103
21:00	ACC70	Accumulated (70 cm) snow	12	95

# MEASUREMENTS, $L^{\uparrow}$

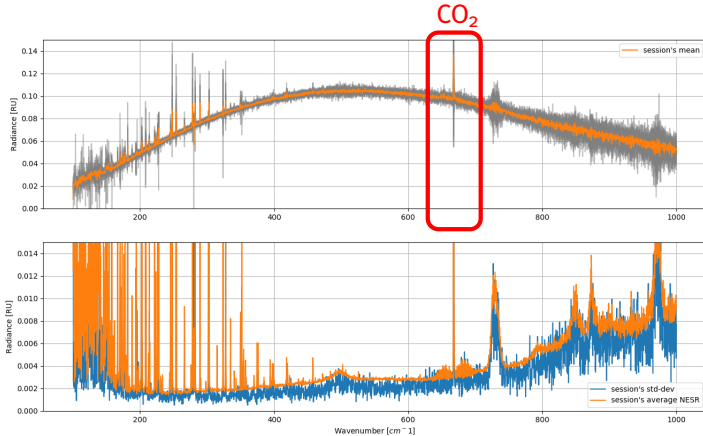




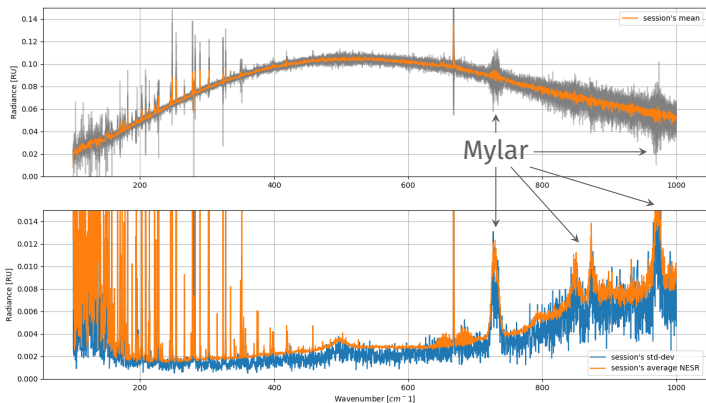
# MEASUREMENTS, $L^\uparrow$



# MEASUREMENTS, $L^{\uparrow}$



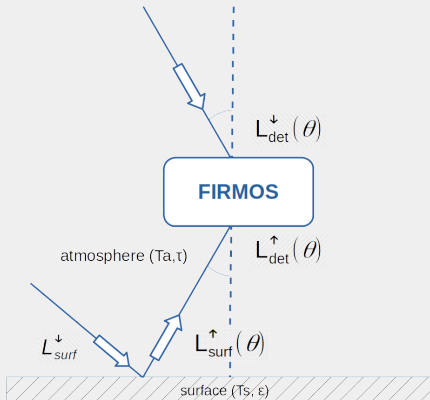
# MEASUREMENTS, $L^\uparrow$



# MODELLING OF THE PROBLEM

$$L_{det}^{\uparrow} = \left[ \epsilon B(T_s) + (1 - \epsilon) L_{surf}^{\downarrow} \right]$$

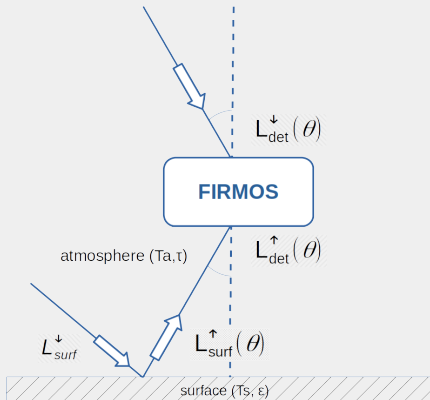
- $\epsilon_{\nu}$  surface spectral emissivity
- $B_{\nu}(T_s)$  Planck emission at surface temperature ( $T_s$ )
- $L_{\nu, surf}^{\downarrow}$  atmospheric downwelling radiance at the surface



# MODELLING OF THE PROBLEM

$$L_{det}^{\uparrow} = \left[ \epsilon B(T_s) + (1 - \epsilon) L_{surf}^{\downarrow} \right] \tau$$

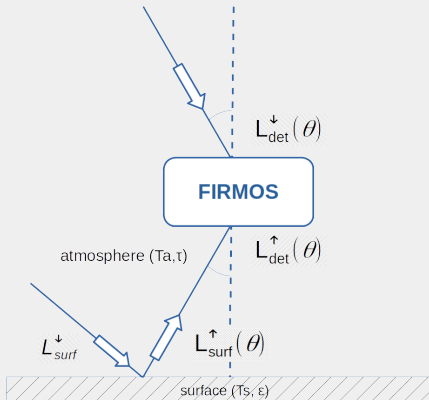
- $\epsilon_{\nu}$  surface spectral emissivity
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- $\tau_{\nu}$  atmospheric spectral transmission



# MODELLING OF THE PROBLEM

$$L_{det}^{\uparrow} = \left[ \epsilon B(T_s) + (1 - \epsilon) L_{surf}^{\downarrow} \right] \tau + E^{\uparrow}$$

- $\epsilon_{\nu}$  surface spectral emissivity
- $B_{\nu}(T_s)$  Planck emission at surface temperature ( $T_s$ )
- $L_{\nu, surf}^{\downarrow}$  atmospheric downwelling radiance at the surface
- $\tau_{\nu}$  atmospheric spectral transmission
- $E_{\nu}$  atmospheric path upwelling radiance



## MODELLING OF THE PROBLEM (2)

$$L_{det}^{\uparrow} = \left[ \epsilon B(T_s) + (1 - \epsilon) L_{surf}^{\downarrow} \right] \tau + E^{\uparrow}$$

- Together with the upwelling radiance measured by the detector we need knowledge of the atmospheric downwelling radiance  $L_{surf}^{\downarrow}$
- If the surface is specular reflector we can use the measurement made by the instrument  $L_{surf}^{\downarrow} \simeq L_{det}^{\downarrow}$

## MODELLING OF THE PROBLEM (2)

$$L_{det}^{\uparrow} = \left[ \epsilon B(T_s) + (1 - \epsilon) L_{surf,eff}^{\downarrow} \right] \tau + E^{\uparrow}$$

- Together with the upwelling radiance measured by the detector we need knowledge of the atmospheric downwelling radiance  $L_{surf}^{\downarrow}$
- If the surface is specular reflector we can use the measurement made by the instrument  $L_{surf}^{\downarrow} \simeq L_{det}^{\downarrow}$
- But snow reflection is Lambertian, we model the downwelling radiance at a single effective angle ( $55^\circ$ , Bellisario et al. 2017) using the state of the atmosphere retrieved from the sky observations
- $L_{surf,eff}^{\downarrow} = RT(T, wv, \dots, \theta = 55^\circ)$



# INVERSION

$$L_{det}^{\uparrow} = F \left( \epsilon, T_s, T_a, wv_a, L_{surf,eff}^{\downarrow} \right)$$

We use OE to find an optimal solution  $\hat{x}$  given the observation and prior information using an iterative process

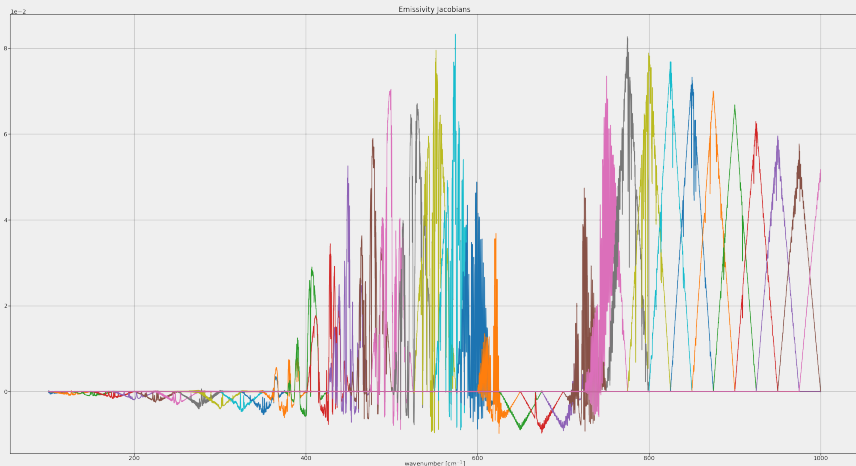
$$x_{i+1} = x_a + (S_a^{-1} + K_i^T S_y^{-1} K_i)^{-1} K_i^T S_y^{-1} [y - F(x_i) + K_i(x_i - x_a)]$$

$$S_i = (S_a^{-1} + K_i^T S_y^{-1} K_i)^{-1}$$

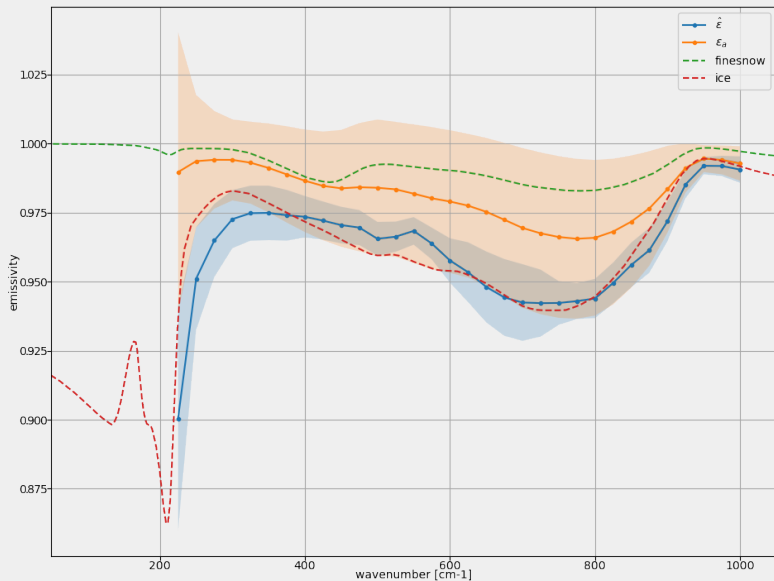
- $x$  state (emissivity, surface T, air humidity and T)
- $x_a, S_a$  prior information
- $K = \frac{\delta y}{\delta x}$  Jacobian matrix to linearise  $F$
- $y, S_y$  observations

# RETRIEVAL

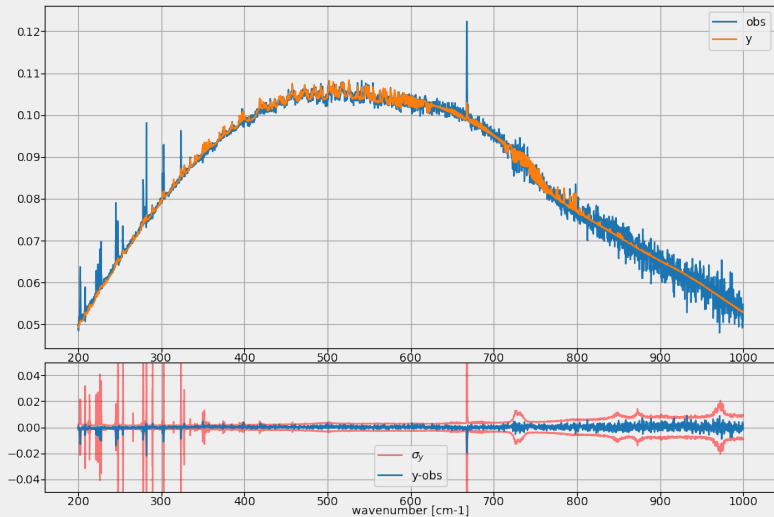
- a-priori - courtesy of X. Huang, same as Huang et al. (2016) but at FIRMOS' viewing angle
- Jacobians



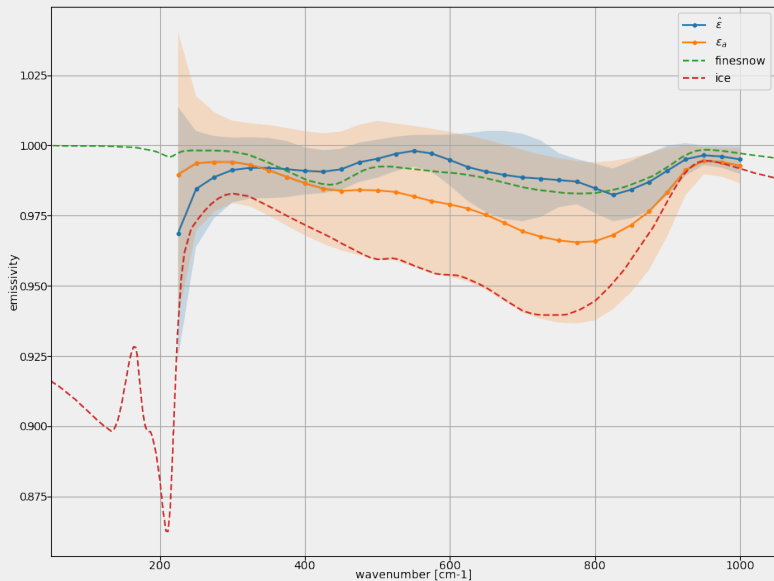
# RESULTS, ICE



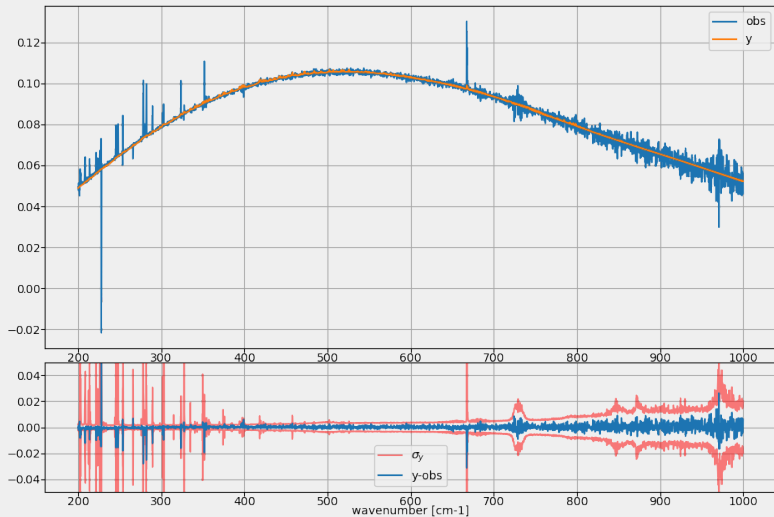
# RESULTS, ICE



# RESULTS, SNOW ACC100



# RESULTS, SNOW ACC100



- The retrieval of snow surface emissivity from FIRMOS shows promising results
- There is good sensitivity in the  $400\text{-}600\text{ cm}^{-1}$  range
- The origin of the dip in the  $200\text{-}400\text{ cm}^{-1}$  range is unclear
- Work is in progress to exploit the snow samples characterisation
- in future campaigns the measurements should alternate between surface and atmosphere

THANK YOU